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Random Error in Space-Time Bin Averages of Sea Surface Temperature Observations from Ships

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Key Points:

- Relative biases between SST data sets from ships and satellites, averaged to one degree monthly bins, are estimated as climatological means
- Magnitudes of difference anomalies between one degree monthly averages of SST from ships and satellites agree with the random error model
- Separate estimates are obtained for sampling and measurement error components of the total error in bin averages of ship SST observations

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Abstract

Sea surface temperature (SST) observations made at ships are distributed irregularly in space and time and are affected by systematic biases and random errors. Such observations are often “binned”: split into samples, contained within “bins” – grid boxes of a space-time grid ($1^\circ \times 1^\circ$ monthly bins are used here), and their statistics are computed. Bin averages often serve as gridded representations of such data, thus requiring reliable uncertainty estimates, which for ship observations are particularly important because of their domination in the early observational records. Here ship SST observations for 1992–2010 are compared with an independent high-resolution satellite-based SST data set. To remove systematic biases, seasonal means were subtracted from the difference between bin-averaged data sets. In more than 66%(50%) of locations with binned temporal coverage exceeding 50%(66%), the magnitude of remaining anomalies agreed within 20%(10%) with random error model estimates. Separate estimates for sampling and measurement error components were obtained.

Plain Language Summary

Sea surface temperature (SST) is an important climate variable. SST observations made at ships are distributed irregularly in space and time and are affected both by systematic biases and randomly-varying measurement errors. To make them easier to use, such data sets are often “binned”, i.e., split into samples contained within “bins”, which usually are grid boxes of some space-time grid (monthly 1° longitude by 1° latitude bins are used here), and the statistics of these binned samples are computed. Bin averages often serve as gridded representations for data sets of ship observations; hence their uncertainty estimates have to be reliable. This is especially important since ship observations dominate early on in the historical observational record. Ship SST observations for 1992–2010 are compared here with an independent high-resolution satellite-based SST data set. To remove systematic biases, seasonal means were subtracted from the difference between bin-averaged versions of these data sets, and the remainder was interpreted as a sum of random errors. In more than 66%(50%) of locations where binned temporal coverage exceeded 50%(66%), the remainder’s magnitude agreed within 20%(10%) with the random error model estimates. Error components due to incomplete sampling and due to measurement error were estimated.

1 Introduction

Sea surface temperature (SST) is one of the “essential” climate variables (Bojinski et al., 2014), particularly well-suited for monitoring changes in the Earth’s mean surface temperature and very visible in the climate change debate (Hartmann et al., 2013). More than two centuries of SST observations together with other in situ data for surface ocean are assembled in the International Comprehensive Ocean-Atmosphere Data Set (ICOADS, Woodruff et al., 1987; Freeman et al., 2017). These observations are irregularly distributed in space and time. A typical preparatory step for their use in climate studies is “binning,” i.e., splitting them into subsamples, contained in non-overlapping spatiotemporal “bins”, usually grid boxes of a regular space-time grid (hereinafter, “bins” are monthly $1^\circ \times 1^\circ$ grid boxes), and reporting statistical summaries of each bin’s sample, e.g., number of observations N_o in the bin, their sample mean \mathcal{M}_o , standard deviation (SD) \mathcal{S}_o , etc. By construction, each of these statistics forms a gridded field, albeit usually incomplete. In lieu of averages over the complete bin’s volume, which are generally unavailable, bin means \mathcal{M}_o (a.k.a. “super-observations”: Smith & Reynolds, 2005; Kennedy, 2014) are often used as input data for objective analyses or data assimilation; hence having reliable uncertainty estimates for binned data averages is important. For ship data this importance is especially high because of ships’ domination, as an observational platform, in the early part of historical data record.

If binned observations were independent and identically distributed (i.i.d.) with variance

$$\sigma_{\mathcal{B}_o}^2 \stackrel{\text{def}}{=} \mathbb{E}S_o^2, \quad (1)$$

the error SD estimate (ESDE) for \mathcal{M}_o would be

$$e_{\mathcal{M}_o} \stackrel{\text{def}}{=} \sigma_{\mathcal{B}_o} / \sqrt{\mathcal{N}_o}. \quad (2)$$

Hereinafter label “def” above the “=” sign introduces an expression on the left of this sign as denoting the expression on its right, and \mathbb{E} denotes mathematical expectation. Error estimates computed by (2), but with $\sigma_{\mathcal{B}_o}$ calculated as the root-mean-square (RMS) over the best data coverage period of all bins’ SDs in that location, was introduced by Kaplan et al. (1997) and used for objective analyses of historical SST observations by Kaplan et al. (1998), Ilin and Kaplan (2009), and, with further modifications to $\sigma_{\mathcal{B}_o}$ estimate, by Karspeck et al. (2012). Their analyzed fields and uncertainty estimates provide some indirect validation for such uses of formula (2).

However, a direct comparison of the RMS difference between ICOADS bin means and satellite SST data with error estimates based on (2), while showing general large-scale agreement between global patterns of error magnitude, had many regional and smaller-scale differences (Rayner et al., 2010, cf. their Figures 1e vs. 1f). There were many possible reasons for this lack of detailed agreement: likely failure of the i.i.d. assumption, since the binned samples included observations from different platform types (ships, moorings, drifting buoys), obtained by different observational methods and affected by biases associated with particular methods of observations, platform types, or individual platforms (e.g., persistent thermometer biases on individual ships). Furthermore, the interpretation of that comparison was complicated by the dependence of the satellite-based SST data set on the in situ observations themselves.

A high-resolution interpolated SST analysis product (available on a daily $0.05^\circ \times 0.05^\circ$ grid), based on the satellite data, independent of the concurrent in situ SST observations, and accompanied by verified uncertainty estimates, had become available several years ago (Merchant et al., 2014). The actual RMS differences for 1992-2010 between bin-averaged SST from ship observations and from this independent satellite-based analysis are compared here with their estimates based on the random error model and the analysis uncertainty estimates. While the bias structure of ship SST observations is complicated and remains a subject of active research (Kent et al., 2017), it is hypothesized here that a large part of biases in bin-averaged ship SST data can be approximated by its seasonally-dependent component. Once the climatological average is removed from the difference between ship and satellite bin means, the residual anomaly is treated as a combination of random errors. Separate estimates are obtained for sampling and measurement error components in bin-averaged ship SST data.

Section 2 describes data sets used and their pre-processing for this study. Section 3 presents error models, their estimates, and the technique of their comparison with the RMS of difference anomaly between bin-averaged versions of ICOADS ship SST and the satellite analysis product. Section 4 presents the results, which are discussed in section 5. Conclusions are given in section 6.

2 Data

2.1 High-resolution satellite SST analysis product

High-resolution globally-complete satellite SST data set, independent of in situ data (Merchant et al., 2014), produced within the Climate Change Initiative (CCI) of the European Space Agency (ESA), is used. It is based on the consistent re-processing of major global streams of the infrared satellite SST data, namely, the data from (Advanced)

110 Along-Track Scanning Radiometer and from the Advanced Very High Resolution Ra-
 111 diometer missions, with the deliberate avoidance of product dependency on the concu-
 112 rent in situ SST observations (coefficients in SST retrievals were computed by optimal
 113 estimation, based on the atmospheric radiative transfer simulations, rather than by best-
 114 fitting in situ SST observations). In addition to the more traditional “skin” SST, the time-
 115 adjusted temperature at 20 cm depth was also produced, by modeling the near-surface
 116 thermally-stratified ocean layer. These temperature values with their uncertainty esti-
 117 mates were fed into the optimal interpolation system for the U.K. Met Office Ocean Sea
 118 Surface Temperature and Sea Ice Analysis (OSTIA, Donlon et al., 2012; Roberts-Jones
 119 et al., 2012, 2016), producing globally-complete ocean temperature fields at 20 cm depth,
 120 $0.05^\circ \approx 6$ km spatial resolution, interpretable as local-time daily averages, with uncertainty
 121 estimates, for 09/1991–12/2010. This product, known as ESA SST CCI Analysis, ver-
 122 sion 1.0, hereinafter will be referred to as “CCI Analysis” or simply “CCI.” The period
 123 of complete 19 years (1992–2010) and $75^\circ\text{S}–75^\circ\text{N}$ global ocean domain will be used.

124 2.2 Ship Observations of SST

125 Ship observations of SST in ICOADS (Release 3.0; Freeman et al., 2017) were iden-
 126 tified by the “Platform Type” indicator value (PT=5), corresponding to the “ship” ob-
 127 servational platform type, and put through the ICOADS own quality control (QC) sys-
 128 tem with settings for the “enhanced Monthly Summary Trimmed Group”. For each ship
 129 SST observation o that passed QC, its local time and date were computed and included
 130 into its record (only Coordinated Universal Time and date are in the ICOADS own data
 131 format). Then “match-up” SST value a^o and its uncertainty SD e^{a^o} from the CCI Anal-
 132 ysis to this ship observation o were extracted and also added to the record for o . Match-
 133 up values from the CCI Analysis to the given ship observation are the analyzed SST value
 134 and its uncertainty SD for the daily $0.05^\circ \times 0.05^\circ$ grid box within whose time-space lim-
 135 its this ship observation was made. (A small number of ICOADS ship SST observations
 136 that passed QC but did not have CCI Analysis match-up values were excluded from this
 137 study.)

138 2.3 Data Preparation

139 Consider bin \mathcal{B} , representing a grid box of a regular monthly $1^\circ \times 1^\circ$ grid, and a sam-
 140 ple \mathcal{B}_o of \mathcal{N}_o SST observations from ships that were taken within its space and time lim-
 141 its and successfully passed ICOADS QC:

$$142 \mathcal{B}_o \stackrel{\text{def}}{=} \{o_1, o_2, \dots, o_{\mathcal{N}_o}\}.$$

143 This “binned” sample is characterized by its mean \mathcal{M}_o and SD \mathcal{S}_o , as follows:

$$144 \mathcal{M}_o \stackrel{\text{def}}{=} \frac{1}{\mathcal{N}_o} \sum_{i=1}^{\mathcal{N}_o} o_i, \quad \mathcal{S}_o^2 \stackrel{\text{def}}{=} \frac{1}{\mathcal{N}_o - 1} \sum_{i=1}^{\mathcal{N}_o} (o_i - \mathcal{M}_o)^2. \quad (3)$$

145 Note that \mathcal{S}_o can only be computed if $\mathcal{N}_o > 1$. Therefore bins with only one ship obser-
 146 vation ($\mathcal{N}_o = 1$) form a special class of data samples: their means, but not variability can
 147 be estimated directly from their data. Dealing with this more complicated subset is left
 148 for further investigation, and only bins with $\mathcal{N}_o \geq 2$ are considered in this study.

149 Consider also a set \mathcal{B}_a of \mathcal{N}_a SST values from the CCI Analysis for all daily $0.05^\circ \times 0.05^\circ$
 150 grid boxes contained within that same bin \mathcal{B} as above:

$$151 \mathcal{B}_a \stackrel{\text{def}}{=} \{a_1, a_2, \dots, a_{\mathcal{N}_a}\}.$$

152 Its statistics \mathcal{M}_a and \mathcal{S}_a are computed similarly to (3):

$$153 \mathcal{M}_a \stackrel{\text{def}}{=} \frac{1}{\mathcal{N}_a} \sum_{j=1}^{\mathcal{N}_a} a_j, \quad \mathcal{S}_a^2 \stackrel{\text{def}}{=} \frac{1}{\mathcal{N}_a - 1} \sum_{j=1}^{\mathcal{N}_a} (a_j - \mathcal{M}_a)^2. \quad (4)$$

154 Unless the land or ice cover are present within the bin \mathcal{B} , the number of data points in
 155 \mathcal{B}_a is quite large: typically, $\mathcal{N}_a \sim 20 \times 20 \times 30 \gg \mathcal{N}_o$ for ocean locations. In fact, con-
 156 straint $\mathcal{N}_a \geq 2$ is satisfied automatically for all non-empty binned samples \mathcal{B}_a from the
 157 CCI Analysis.

158 Recall that for each $o_i \in \mathcal{B}_o$, its CCI SST match-up a_i^o has been identified and
 159 stored in the record for o_i (Section 2.2). Therefore it is easy to assemble a sample of CCI
 160 Analysis match-ups to ship observations in \mathcal{B}_o :

$$161 \quad \mathcal{B}_{ao} \stackrel{\text{def}}{=} \{a_1^o, a_2^o, \dots, a_{\mathcal{N}_o}^o\}$$

162 and to compute its statistics \mathcal{M}_{ao} and \mathcal{S}_{ao} analogously to (3). Additionally, differences
 163 between ship observations and their CCI Analysis SST match-ups

$$164 \quad d_i \stackrel{\text{def}}{=} o_i - a_i^o, \quad i = 1, \dots, \mathcal{N}_o \quad (5)$$

165 are binned as well, resulting in the sample

$$166 \quad \mathcal{B}_d \stackrel{\text{def}}{=} \{d_1, d_2, \dots, d_{\mathcal{N}_o}\}$$

167 and its bin statistics \mathcal{M}_d and \mathcal{S}_d .

168 It will prove useful to have bin statistics for CCI Analysis uncertainties pre-computed
 169 as well. These are calculated in exactly the same way as was done above for correspond-
 170 ing SST values. Specifically, let

$$171 \quad \mathcal{B}_{ea} \stackrel{\text{def}}{=} \{e_1^a, e_2^a, \dots, e_{\mathcal{N}_a}^a\},$$

172 where each e_j^a is the uncertainty SD for the CCI Analysis SST value $a_j \in \mathcal{B}_a$ and com-
 173 pute \mathcal{M}_{ea} , \mathcal{S}_{ea} analogously to (4). For the sample of the CCI Analysis uncertainty value
 174 match-ups to the ship observations in \mathcal{B}

$$175 \quad \mathcal{B}_{eao} \stackrel{\text{def}}{=} \{e_1^{ao}, e_2^{ao}, \dots, e_{\mathcal{N}_o}^{ao}\},$$

176 where each e_i^{ao} is the uncertainty SD for the CCI Analysis SST value $a_i^o \in \mathcal{B}_{ao}$ and com-
 177 pute \mathcal{M}_{eao} , \mathcal{S}_{eao} analogously to (3).

178 Described above calculations of \mathcal{M}_x and \mathcal{S}_x are done for all monthly $1^\circ \times 1^\circ$ bins
 179 with sample sizes larger than one, i.e., for bins with $\mathcal{N}_o \geq 2$, when $x = o, ao, d$, or eao ,
 180 and for all bins with non-empty samples when $x = a$ or ea . Temporal attribution of
 181 bin statistics $\mathcal{M}_x(y, m)$, $\mathcal{S}_x(y, m)$ is done using climatological (calendar) month $m =$
 182 $1, \dots, 12$ (January–December) and year $y = 1, \dots, 19$ (corresponding to 1992–2010).
 183 Depending on location, statistics $\mathcal{M}_x(y, m)$, $\mathcal{S}_x(y, m)$ for $x = o, ao, d$, and eao might
 184 not be available for all (y, m) combinations. To simplify the formal treatment of their tem-
 185 poral averaging, for each location if there are $Y_m > 0$ years for month m with $\mathcal{N}_o \geq 2$,
 186 other years are skipped, while available years are renumbered as $y = 1, \dots, Y_m$ for that
 187 month. If only $M < 12$ climatological months are available, these are renumbered as well
 188 from 1 to M . This change in temporal arguments is applied to statistics \mathcal{N}_o and \mathcal{M}_x , \mathcal{S}_x
 189 for $x = o, ao, d$, and eao , as well as for \mathcal{M}_a and \mathcal{M}_{ea} , even though the latter two are
 190 available for all (y, m) for bins in ocean locations; statistics \mathcal{N}_a , \mathcal{S}_a , \mathcal{S}_{ea} remain attributed
 191 to the full set of months and years in 1992–2010 period.

192 With these definitions, available bin averages of SST from ship observations \mathcal{M}_o
 193 and corresponding bin averages from CCI Analysis \mathcal{M}_a have the same temporal argu-
 194 ments, with the timeseries length

$$195 \quad N \stackrel{\text{def}}{=} \sum_{m=1}^M Y_m. \quad (6)$$

196 Hence for their differences

$$197 \quad d_{\mathcal{M}} \stackrel{\text{def}}{=} \mathcal{M}_o - \mathcal{M}_a, \quad (7)$$

198 the full period RMS is computed as

$$199 \quad \mathcal{D} \stackrel{\text{def}}{=} \left[\frac{1}{N} \sum_{m=1}^M \sum_{y=1}^{Y_m} d_{\mathcal{M}}(y, m)^2 \right]^{1/2}. \quad (8)$$

200 3 Methods

201 3.1 Models and assumptions

202 CCI Analysis values a_j are estimates of water temperature at 20 cm depth, averaged
 203 over daily $0.05^\circ \times 0.05^\circ$ grid boxes. Corresponding “true” values t_j^a are averages of
 204 true water temperature t at 20 cm depth over such grid boxes, so for values within bin
 205 \mathcal{B}

$$206 \quad t_j^a = a_j + \varepsilon_j^a, \quad j = 1, \dots, \mathcal{N}_a; \quad (9)$$

$$207 \quad \mathbb{E}\varepsilon_j^a = 0, \quad \mathbb{E}(\varepsilon_j^a)^2 = (e_j^a)^2, \quad j = 1, \dots, \mathcal{N}_a, \quad (10)$$

208 where ε_j^a are the CCI analysis errors. These are assumed uncorrelated with the analyzed
 209 values a_j , since the CCI Analysis is a form of optimal interpolation (Lorenz, 1986). Anal-
 210 ysis errors for different grid boxes, however, are not mutually independent, especially when
 211 these are not greatly separated in time and space. CCI Analysis uses the increased range
 212 (20–350 km) of spatial decorrelation scales of background error that resulted in improved
 213 feature resolution (Roberts-Jones et al., 2016), hence the analysis error is likely domi-
 214 nated by spatial scales larger than $1^\circ \times 1^\circ$. Since the OSTIA background solution uses
 215 day-to-day persistence and relaxes to reference climatology with the 30 day decorrela-
 216 tion time scale (Donlon et al., 2012), a near-perfect correlation of the analysis error within
 217 monthly $1^\circ \times 1^\circ$ bins is assumed here:

$$218 \quad \mathbb{E}(\varepsilon_j^a \varepsilon_k^a) \approx e_j^a e_k^a, \quad j, k = 1, \dots, \mathcal{N}_a. \quad (11)$$

219 For conceptual simplicity, the same “truth” definition, as for the CCI Analysis (9),
 220 is used for ship observations as well:

$$221 \quad o_i = t_i^{ao} + b + \varepsilon_i^o, \quad i = 1, \dots, \mathcal{N}_o, \quad (12)$$

222 where t_i^{ao} is true 20 cm depth temperature averaged over the daily $0.05^\circ \times 0.05^\circ$ grid box
 223 containing ship observation o_i , bias b is assumed constant within each $1^\circ \times 1^\circ$ monthly
 224 bin, thus it does not depend on i in (12). Measurement errors ε_i^o are assumed indepen-
 225 dent of true temperature variations t_i^{ao} and i.i.d. within each bin, with

$$226 \quad \mathbb{E}\varepsilon_i^o = 0, \quad \mathbb{E}(\varepsilon_i^o)^2 = \sigma_o^2, \quad i = 1, \dots, \mathcal{N}_o, \quad (13)$$

227 where σ_o^2 is an (unknown) measurement error variance.

228 Introduce bin average and intra-bin variance of true SST

$$229 \quad \theta \stackrel{\text{def}}{=} \mathcal{M}_{ta}, \quad v^2 \stackrel{\text{def}}{=} \frac{\mathcal{N}_a - 1}{\mathcal{N}_a} \mathcal{S}_{ta}^2, \quad (14)$$

230 where \mathcal{M}_{ta} and \mathcal{S}_{ta} are the mean and variance of the set of true SST values in the bin
 231 \mathcal{B} :

$$232 \quad \mathcal{B}_{ta} \stackrel{\text{def}}{=} \{t_1^a, t_2^a, \dots, t_{\mathcal{N}_a}^a\}.$$

233 Note that v definition above changes the denominator of \mathcal{S}_{ta}^2 from $\mathcal{N}_a - 1$ to \mathcal{N}_a , since (14)
 234 uses discrete analogues of the true value integrals over the bin, rather than statistical
 235 estimates.

Another important assumption is that times and locations of ship observations are random and uniformly distributed over the bin's volume. Hence the true SST match-ups to them form a set of \mathcal{N}_o equiprobable draws

$$\mathcal{B}_{tao} \stackrel{\text{def}}{=} \{t_1^{ao}, t_2^{ao}, \dots, t_{\mathcal{N}_o}^{ao}\}$$

from the full set \mathcal{B}_{ta} of the true SST values in the bin. Based on statistical theorems that lay the foundation of the classical Monte Carlo method for evaluating definite integrals (e.g., Section 3.2 of Robert & Casella, 2004), sample mean \mathcal{M}_{tao} and variance \mathcal{S}_{tao}^2 of this set of random draws \mathcal{B}_{tao} are unbiased estimates of the true mean and variance of the bin

$$\mathbb{E}\mathcal{M}_{tao} = \theta, \quad \mathbb{E}\mathcal{S}_{tao}^2 = v^2, \quad (15)$$

and the error of sample mean, a.k.a. sampling error,

$$\varepsilon_s \stackrel{\text{def}}{=} \mathcal{M}_{tao} - \theta$$

has variance

$$\mathbb{E}\varepsilon_s^2 = v^2/\mathcal{N}_o \quad (16)$$

(for detailed derivation see Section 2.10 of Cochran, 1997).

3.2 Single bin statistics

3.2.1 CCI Analysis samples

Averaging equations (9) over j and using (14), obtain

$$\theta = \mathcal{M}_a + \mathcal{M}_{\varepsilon a}, \quad (17)$$

where $\mathcal{M}_{\varepsilon a}$ is the CCI Analysis error, averaged over the bin. Based on (10) and (11),

$$\mathbb{E}\mathcal{M}_{\varepsilon a} = 0, \quad e_{\mathcal{M}_a}^2 \stackrel{\text{def}}{=} \mathbb{E}\mathcal{M}_{\varepsilon a}^2 = \frac{1}{\mathcal{N}_a^2} \sum_{j,k=1}^{\mathcal{N}_a} \mathbb{E}(\varepsilon_j^a \varepsilon_k^a) \approx \frac{1}{\mathcal{N}_a^2} \sum_{j,k=1}^{\mathcal{N}_a} e_j^a e_k^a = \frac{1}{\mathcal{N}_a^2} \left(\sum_{j=1}^{\mathcal{N}_a} e_j^a \right)^2 = \mathcal{M}_{\varepsilon a}^2. \quad (18)$$

Subtracting (17) from (9), averaging squares of both sides over j , find, using (14), for mathematical expectation of both sides

$$\frac{\mathcal{N}_a}{\mathcal{N}_a - 1} v^2 = \mathbb{E}\mathcal{S}_a^2 + \mathbb{E}\mathcal{S}_{\varepsilon a}^2, \quad (19)$$

where $\mathcal{S}_{\varepsilon a}^2$ is the sample variance of the CCI analysis error in the bin. Using (10), (11), and (18), derive

$$\mathbb{E}\mathcal{S}_{\varepsilon a}^2 = \frac{1}{\mathcal{N}_a - 1} \sum_{i=1}^{\mathcal{N}_o} \mathbb{E}(\varepsilon_i^a)^2 - \frac{\mathcal{N}_a}{\mathcal{N}_a - 1} \mathbb{E}\mathcal{M}_{\varepsilon a}^2 = \mathcal{S}_{\varepsilon a}^2. \quad (20)$$

Substituting (20) into (19), find that

$$\hat{v}^2 = \frac{\mathcal{N}_a - 1}{\mathcal{N}_a} (\mathcal{S}_a^2 + \mathcal{S}_{\varepsilon a}^2), \quad (21)$$

is an unbiased estimator of v^2 .

Subselecting from equations (9) those for the CCI analysis match-ups a^o to ship observations from \mathcal{B}_o ,

$$t_i^{ao} = a_i^o + \varepsilon_i^{ao}, \quad i = 1, \dots, \mathcal{N}_o, \quad (22)$$

analogously to the derivation of (21), find that

$$\hat{v}_o^2 = \mathcal{S}_{ao}^2 + \mathcal{S}_{\varepsilon ao}^2, \quad (23)$$

is another unbiased estimator of v^2 , but based on a small subset of CCI Analysis grid points (\mathcal{N}_o match-ups to ship observations) than the full set of \mathcal{N}_a CCI Analysis points in \mathcal{B} , used in (21).

275 **3.2.2 Ship observations sample**

276 Averaging both sides of (12) over i , obtain

277
$$\mathcal{M}_o = \mathcal{M}_{tao} + b + \mathcal{M}_{\varepsilon_o}, \quad (24)$$

278 where $\mathcal{M}_{\varepsilon_o}$ is the bin mean of measurement errors with, based on (13),

279
$$\mathbb{E}\mathcal{M}_{\varepsilon_o}^2 = \sigma_o^2/\mathcal{N}_o. \quad (25)$$

280 Subtracting (24) from (12), averaging squares of both sides over i , find for mathematical expectation of both sides, using (13), (15), and recalling (1),

282
$$\sigma_{\mathcal{B}_o}^2 = \mathbb{E}\mathcal{S}_{tao}^2 + \mathbb{E}\mathcal{S}_{\varepsilon_o}^2 = v^2 + \sigma_o^2. \quad (26)$$

283 Inserting t_i^{ao} from (22) into (12), obtain, recalling (5),

284
$$d_i = b + \varepsilon_i^o + \varepsilon_i^{ao}, \quad i = 1, \dots, \mathcal{N}_o. \quad (27)$$

285 By taking sample variances of both sides of (27) and considering their expectations, find an unbiased estimate of ship SST measurement error σ_o^2 :

287
$$\hat{\sigma}_o^2 = \mathcal{S}_d^2 - \mathcal{S}_{\varepsilon_{ao}}^2. \quad (28)$$

288 **3.2.3 Bin mean differences**

289 For differences between bin-averaged ship observations and CCI Analysis, defined by (5):

291
$$d_{\mathcal{M}} = b + \varepsilon_{d\mathcal{M}}, \quad (29)$$

292 where

293
$$\varepsilon_{d\mathcal{M}} \stackrel{\text{def}}{=} \varepsilon_s + \mathcal{M}_{\varepsilon_o} + \mathcal{M}_{\varepsilon_a},$$

294 and based on (16), (18), (25), and (26),

295
$$\mathbb{E}\varepsilon_{d\mathcal{M}} = 0, \quad \mathbb{E}\varepsilon_{d\mathcal{M}}^2 \stackrel{\text{def}}{=} \mathbb{E}\varepsilon_{d\mathcal{M}}^2 = \sigma_{\mathcal{B}_o}^2/\mathcal{N}_o + \mathcal{M}_{\varepsilon_a}^2. \quad (30)$$

296 **3.3 Statistics for a temporal sample of bins**

297 **3.3.1 Actual RMS differences**

298 Consider a temporal sample of bin statistics for a certain location of the bin. Due to (29), straight RMS \mathcal{D} of differences $d_{\mathcal{M}}(y, m)$, calculated by (8), is affected by bias b . Bias estimate $\hat{b}_c(m)$ is obtained by climatological averaging of $d_{\mathcal{M}}(y, m)$:

301
$$\hat{b}_c(m) = \frac{1}{Y_m} \sum_{y=1}^{Y_m} d_{\mathcal{M}}(y, m), \quad m = 1, \dots, M. \quad (31)$$

302 The RMS of the differences $d_{\mathcal{M}}$ with the estimated bias removed, taking into account the reduction in the number of degrees of freedom (DOF) from (6) to

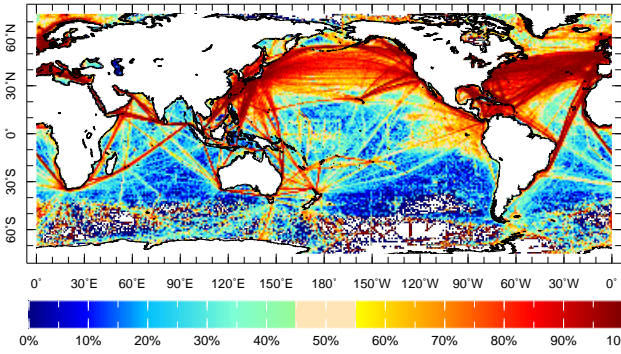
304
$$\sum_{m=1}^M (Y_m - 1) = N - M,$$

305 becomes

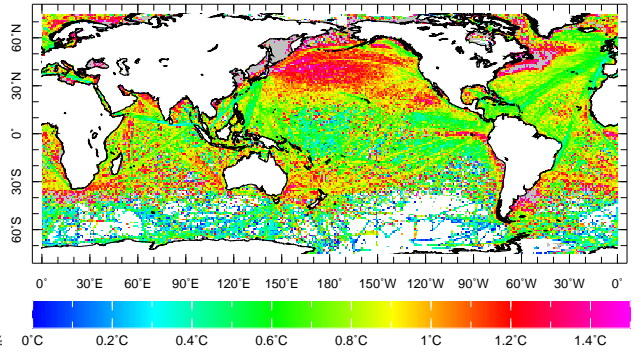
306
$$\mathcal{D}' = \left[\frac{1}{N - M} \sum_{m=1}^M \sum_{y=1}^{Y_m} \left(d_{\mathcal{M}}(y, m) - \hat{b}_c(m) \right)^2 \right]^{1/2}. \quad (32)$$

307

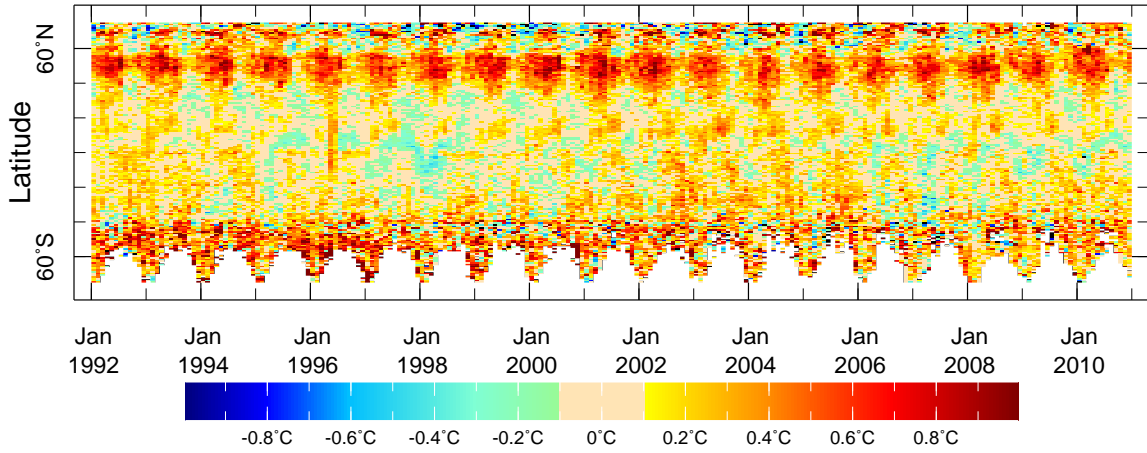
(a) % of monthly $1^\circ \times 1^\circ$ SST bins with $N_o \geq 2$, 47.8%



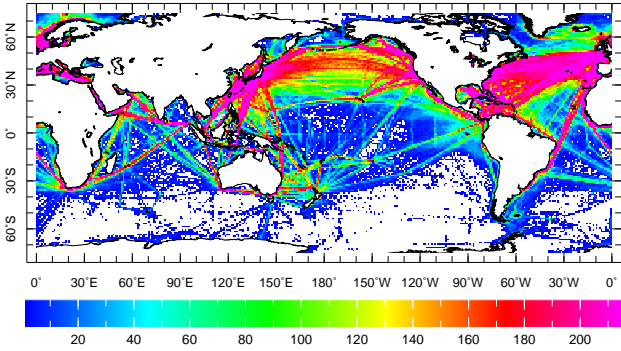
(b) RMS \mathcal{D} of ship-satellite difference d_M , 0.99°C



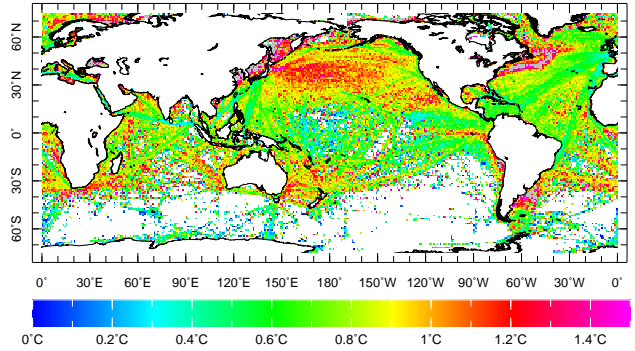
(c) Zonal means of SST differences d_M , $^\circ\text{C}$, between bin-averaged ship observations and CCI Analysis



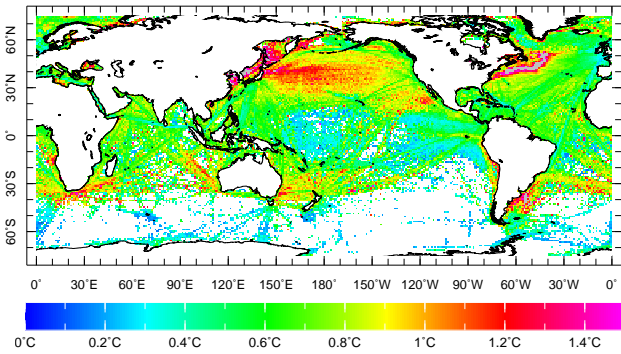
(d) Number of DOF in anomalies, 71.3



(e) RMS \mathcal{D}' of ship-satellite anomaly $d_M - \hat{b}_c$, 0.91°C



(f) Estimate \mathcal{E} of binned anomaly RMS \mathcal{D}' , 0.74°C



(g) Relative discrepancy ρ , 81%

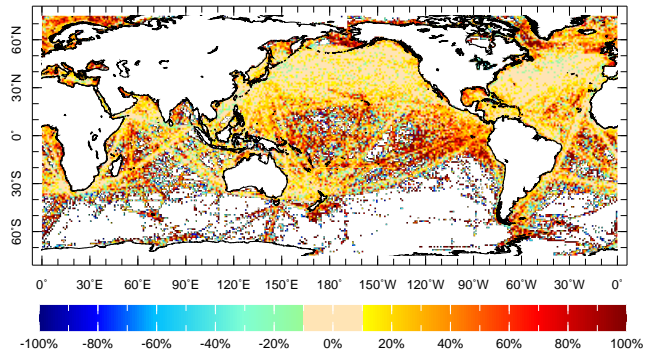


Figure 1. Comparison of monthly $1^\circ \times 1^\circ$ bin-averaged ($\mathcal{N}_o \geq 2$) ICOADS ship SST observations with the ESA CCI Analysis for 1992-2010: (a) Percentage of ICOADS ship SST bins with $\mathcal{N}_o \geq 2$ among all bins with data ($\mathcal{N}_o \geq 1$); (b) RMS \mathcal{D} , $^\circ\text{C}$, of difference $d_{\mathcal{M}}$ between bin-averaged ship and satellite data; (c) Zonal averages of differences $d_{\mathcal{M}}$, $^\circ\text{C}$; (d) DOF in anomalies of bin-averaged ship data (zero DOF grids are shown as missing data, in white); (e) RMS \mathcal{D}' , $^\circ\text{C}$, of ship-satellite difference anomalies $d_{\mathcal{M}} - \hat{b}_c$; (f) Estimate \mathcal{E} , $^\circ\text{C}$, of ship-satellite difference anomaly RMS \mathcal{D}' ; (g) Relative difference ρ , %, between \mathcal{D}' and \mathcal{E} . Numbers at the end of panel labels are: for (a),(d) – global averages of displayed fields; for (b),(e),(f),(g) – global RMS of displayed fields.

3.3.2 Estimated RMS differences and errors

Based on (29) and (31),

$$\begin{aligned} \mathbb{E}\mathcal{D}'^2 &= \frac{1}{N-M} \mathbb{E} \left[\sum_{m=1}^M \sum_{y=1}^{Y_m} \left(\varepsilon_{d\mathcal{M}}(y, m) - \frac{1}{Y_m} \sum_{q=1}^{Y_m} \varepsilon_{d\mathcal{M}}(y, m) \right)^2 \right] = \\ &= \frac{1}{N-M} \sum_{m=1}^M \left[\mathbb{E} \sum_{y=1}^{Y_m} \varepsilon_{d\mathcal{M}}(y, m)^2 - \frac{1}{Y_m} \mathbb{E} \left(\sum_{q=1}^{Y_m} \varepsilon_{d\mathcal{M}}(y, m) \right)^2 \right] = \\ &= \sum_{m=1}^M \frac{\mu_m}{Y_m} \sum_{y=1}^{Y_m} e_{d\mathcal{M}}(y, m)^2, \end{aligned}$$

where

$$\mu_m \stackrel{\text{def}}{=} (Y_m - 1)/(N - M), \quad m = 1, \dots, M. \quad (33)$$

is the portion of the total DOF due to each climatological month m (note that $\sum_{m=1}^M \mu_m = 1$). Based on (30),

$$e_{d\mathcal{M}}(y, m)^2 = \sigma_{\mathcal{B}o}(m)^2 / \mathcal{N}_o(y, m) + \mathcal{M}_{ea}(y, m)^2$$

and

$$\mathbb{E}\mathcal{D}'^2 = \sum_{m=1}^M \mu_m \sigma_{\mathcal{B}o}(m)^2 / \mathcal{N}_o^h(m) + \sum_{m=1}^M \mu_m \mathcal{M}_{ea}^q(m)^2, \quad (34)$$

where

$$\mathcal{N}_o^h(m) \stackrel{\text{def}}{=} \left[\frac{1}{Y_m} \sum_{y=1}^{Y_m} \mathcal{N}_o(y, m)^{-1} \right]^{-1}, \quad \mathcal{M}_{ea}^q(m) \stackrel{\text{def}}{=} \left[\frac{1}{Y_m} \sum_{y=1}^{Y_m} \mathcal{M}_{ea}(y, m)^2 \right]^{1/2}.$$

are harmonic $\mathcal{N}_o(y, m)$ and quadratic $\mathcal{M}_{ea}(m)$ means of \mathcal{N}_o and \mathcal{M}_{ea} , respectively, over available years $y = 1, \dots, Y_m$.

An estimate of $\sigma_{\mathcal{B}o}(m)^2$ is computed as pooled variance (Section 9.2.16 in Von Storch & Zwiers, 2001) of binned samples over all available years $y = 1, \dots, Y_m$:

$$\hat{\sigma}_{\mathcal{B}o}^2(m) \stackrel{\text{def}}{=} \sum_{y=1}^{Y_m} \gamma(y, m) \mathcal{S}_o(y, m)^2, \quad m = 1, \dots, M, \quad (35)$$

where weighting coefficients are

$$\gamma(y, m) \stackrel{\text{def}}{=} [\mathcal{N}_o(y, m) - 1] / \Gamma(m), \quad y = 1, \dots, Y_m, \quad m = 1, \dots, M, \quad (36)$$

$$\Gamma(m) \stackrel{\text{def}}{=} \sum_{y=1}^{Y_m} [\mathcal{N}_o(y, m) - 1], \quad m = 1, \dots, M. \quad (37)$$

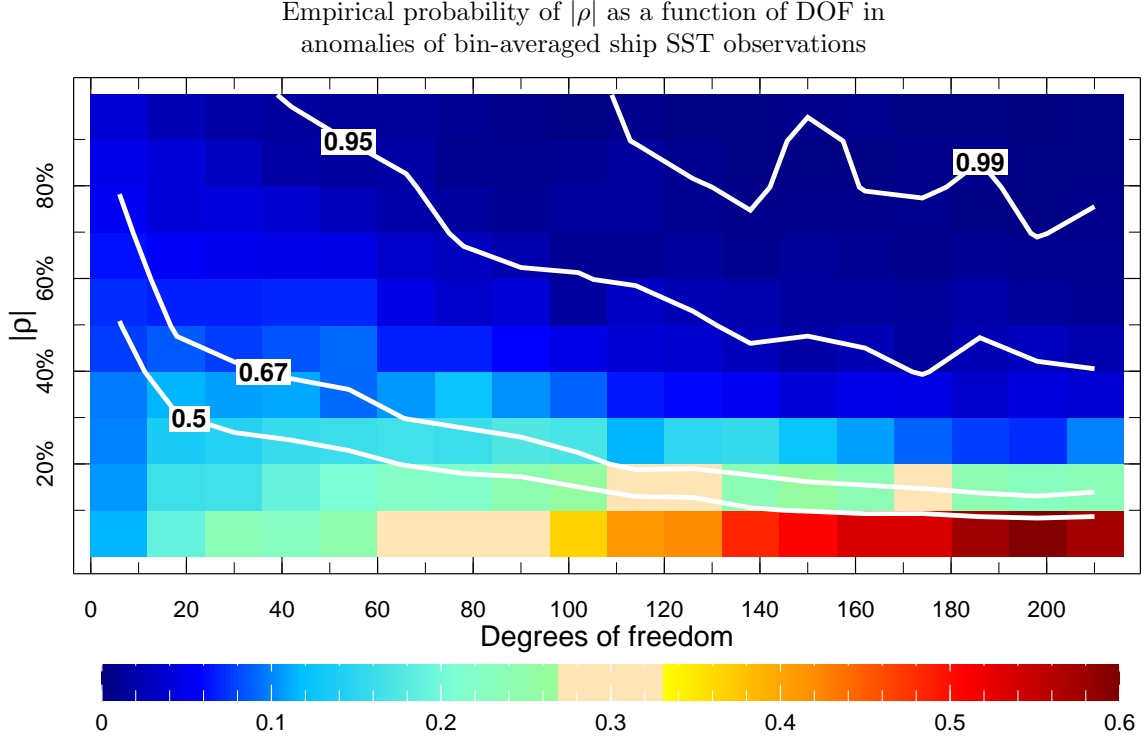


Figure 2. Empirical probability (color) of $|\rho|$ calculated for 10%-wide segments of 0–100% interval (vertical axis) for each of 12-wide sub-ranges of the complete 1–216 range of the possible DOF in the climatological anomaly sample for 1992–2010 (horizontal axis). White lines are contours of cumulative empirical probability of $|\rho|$, conditional on the given DOF range, corresponding to the values of 0.5, 0.67, 0.95, and 0.99, as labels indicate.

Substituting estimate $\hat{\sigma}_{Bo}(m)^2$ from (35) for the value of $\sigma_{Bo}(m)^2$ in (34), obtain an unbiased estimate for \mathcal{D}'^2 :

$$\mathcal{E}^2 \stackrel{\text{def}}{=} \mathcal{E}_{\mathcal{M}o}^2 + \mathcal{E}_{\mathcal{M}a}^2, \quad (38)$$

where the terms in the right-hand side are estimates of error variances in bin averages of ship observations

$$\mathcal{E}_{\mathcal{M}o}^2 \stackrel{\text{def}}{=} \sum_{m=1}^M \mu_m \hat{\sigma}_{Bo}(m)^2 / \mathcal{N}_o^h(m) \quad (39)$$

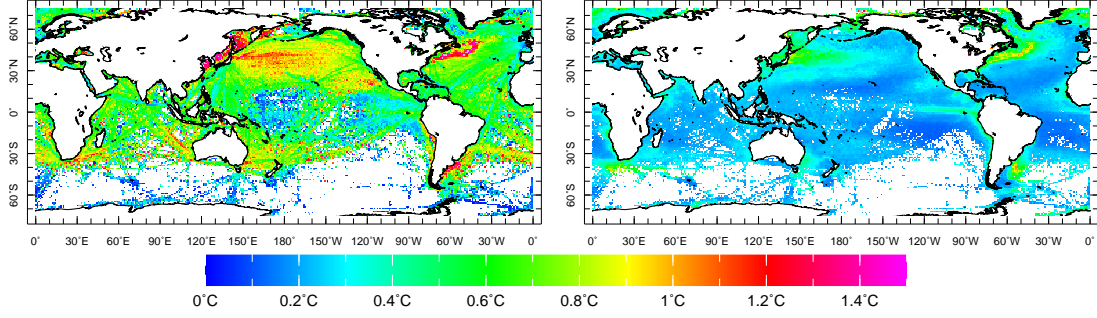
and of the CCI Analysis

$$\mathcal{E}_{\mathcal{M}a}^2 \stackrel{\text{def}}{=} \sum_{m=1}^M \mu_m \mathcal{M}_{ca}^q(m)^2. \quad (40)$$

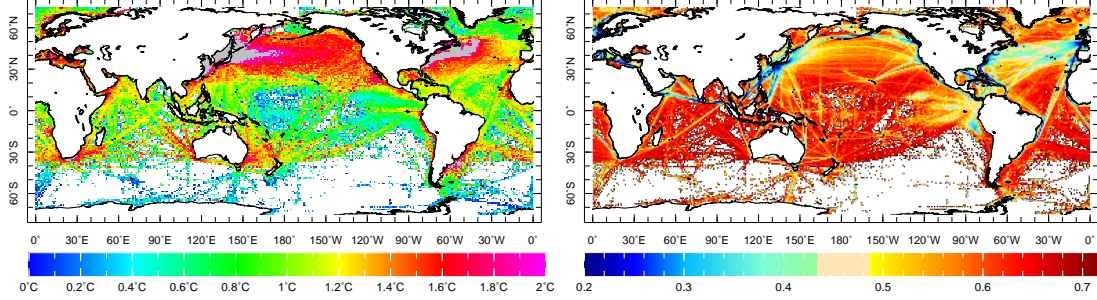
4 Results

Excluded from this study monthly $1^\circ \times 1^\circ$ bins with a single ship SST observation constitute a surprisingly large percentage (31.8%) of all ICOADS 1992–2010 monthly $1^\circ \times 1^\circ$ bins with ship SST observations (with any $\mathcal{N}_o > 0$). Figure 1a shows local percentages of bins included in this study ($\mathcal{N}_o \geq 2$) among all bins with ship SST data ($\mathcal{N}_o > 0$), identifying better-sampled areas in North Atlantic and North Pacific Oceans and along ship tracks. Figure 1b shows RMS \mathcal{D} of differences $d_{\mathcal{M}}$ between bin averages of ship SST observations and CCI Analysis for 1992–2010 (see equations (7) and (8)).

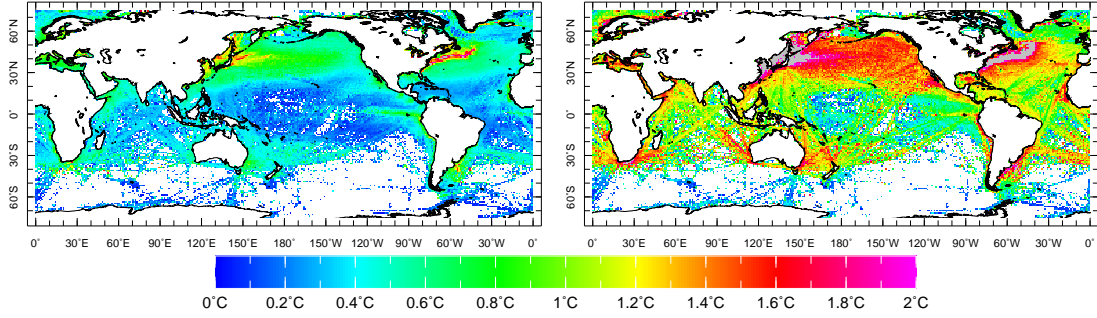
(a)ERMSE \mathcal{E}_{Mo} of bin-averaged ship SST, 0.67°C (b)ERMSE \mathcal{E}_{Ma} of bin-averaged CCI, 0.31°C



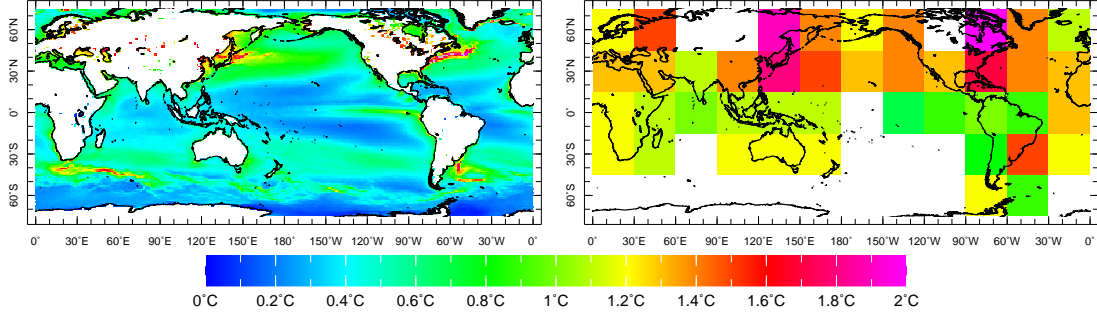
(c)Intra-bin SD of ship observations $\hat{\sigma}_{Bo}^*$, 1.20°C (d) Error reduction factor $1/\sqrt{N^*}$, 0.57



(e) Sampling ESDE $\hat{\sigma}_o^*$, 0.50°C (f) Ship measurement ESDE $\hat{\sigma}_o^*$, 1.14°C



(g) Sampling ESDE $\hat{\sigma}_o^*$, 0.57°C (h) Kent&Challenor(2006) ship ESDE, 1.26°C



(i) same as (f), but in $30^\circ \times 30^\circ$ averages, 1.13°C (j) Relative difference ρ : (h) vs (i), 18%

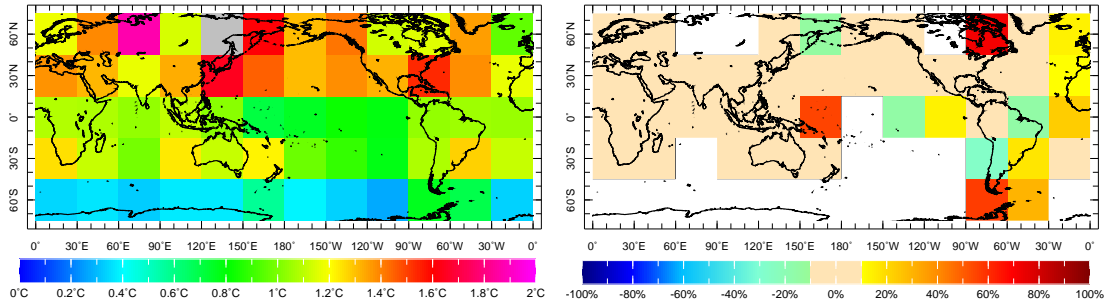


Figure 3. Components of estimated RMS difference anomaly between bin-averaged ship SST observations and CCI Analysis: (a) ERMSE $\mathcal{E}_{\mathcal{M}_o}$ of bin-averaged ship SST, °C; (b) ERMSE $\mathcal{E}_{\mathcal{M}_a}$ of bin-averaged CCI Analysis SST, °C; (c) Intra-bin SD $\hat{\sigma}_{\mathcal{B}_o}^*$ of ship observations, °C; (d) Average error reduction factor $1/\sqrt{\mathcal{N}^*}$; (e) Sampling ESDE \hat{v}_o^* from the CCI Analysis match-ups to ship observations, °C; (f) Measurement ESDE $\hat{\sigma}_o^*$ of ship SST, °C; (g) Sampling ESDE \hat{v}^* from the full CCI Analysis, °C; (h) ship SST random ESDE from Kent and Challenor (2006, their Figure 2), °C; (i) same as (f), but in $30^\circ \times 30^\circ$ averages, °C; (j) Relative difference ρ between (h) and (i), %. Numbers at the end of panel labels indicate displayed fields' global RMS.

346 These differences have substantial mean and seasonal components, as seen in the
 347 time-latitude plot of zonally-averaged $d_{\mathcal{M}}$ (Figure 1c). Subtracting from $d_{\mathcal{M}}$ their cli-
 348 matological mean reduces the DOF by one for each climatological month, represented
 349 in the data (Figure 1d), but even accounting for the reduced DOF, the RMS \mathcal{D}' of $d_{\mathcal{M}}$
 350 anomaly, calculated by (32) and shown in Figure 1e, is appreciably smaller than \mathcal{D} (8%
 351 global RMS reduction).

352 The difference $d_{\mathcal{M}}$ anomaly is interpreted here as the sum of random errors in bin
 353 averages of ship observations and of CCI analysis; the estimate \mathcal{E} of its RMS \mathcal{D}' , based
 354 on this model, is computed by equation (38) and shown in Figure 1f. It matches \mathcal{D}' pat-
 355 tern (Figure 1e) in many details. To aid their visual comparison, their difference

$$\rho = (\mathcal{D}' - \mathcal{E}) / \mathcal{E}$$

357 is expressed as the percentage of the estimate \mathcal{E} and is shown in Figure 1g, where large
 358 areas of the actual and estimated RMS agreeing within 10% or so are clearly seen.

359 The areas of poor agreement in Figure 1g appear to collocate with areas of smaller
 360 DOF in Figure 1d. To quantify this relationship, the empirical probability of $|\rho|$ in 10%
 361 intervals is shown in Figure 2 for different 12-wide DOF ranges of $d_{\mathcal{M}}$ anomalies (1-12,
 362 13-24,..., 205-216). As DOF increases, $|\rho|$ concentrates more in its interval of smallest
 363 values. For more than 67% of points where DOF exceeds 50% of its maximum value (108=
 364 $0.5 \times 12 \times (19-1)$), $|\rho| < 20\%$; for more than half of the points, where for DOF exceeds 144
 365 ($2/3$ of its maximum), $|\rho| < 10\%$.

366 The variance of difference anomaly between bin-averaged ship SST and CCI Analysis
 367 is modeled by (38), as a sum of squares of two components: estimated RMS error
 368 (ERMSE) $\mathcal{E}_{\mathcal{M}_o}$ of bin-averaged ship observations, calculated by (39) and shown in Fig-
 369 ure 3a, and ERMSE of bin-averaged CCI Analysis $\mathcal{E}_{\mathcal{M}_a}$, calculated using (40) and shown
 370 in Figure 3b. The former clearly dominates: the CCI analysis error represents only 17.6%
 371 of the global variance in the total ERMSE \mathcal{E} (Figure 1f).

372 As seen from (39), ERMSE for bin-averaged ship observations averages over the
 373 climatological month m products of intra-bin variance estimates $\hat{\sigma}_{\mathcal{B}_o}(m)^2$ with inverse
 374 harmonic means $1/\mathcal{N}_o^h(m)$ of observational counts. Figures 3c,d show square roots of these
 375 quantities averaged over available climatological months:

$$\hat{\sigma}_{\mathcal{B}_o}^* \stackrel{\text{def}}{=} \left[\sum_{m=1}^M \mu_m \hat{\sigma}_{\mathcal{B}_o}^2(m) \right]^{1/2}, \quad 1/\sqrt{\mathcal{N}^*} \stackrel{\text{def}}{=} \left[\sum_{m=1}^M \mu_m / \mathcal{N}_o^h(m) \right]^{1/2}, \quad (41)$$

377 where μ_m are defined by (33).

378 Figure 3c shows, in effect, the ESDE for the bin-averaged ship SST, if all monthly
 379 bins in the given location only had single observations in them, while 3d shows the ESDE
 380 reduction factor due to the multiple observations. Because of $\mathcal{N}_o \geq 2$ constraint, all val-
 381 ues shown in 3d do not exceed $\sqrt{1/2} \approx 0.71$; their global RMS is 0.57, and the reduc-
 382 tions to much smaller factors are relatively rare: the interquartile range is 0.51–0.64, and
 383 only 3.1% of shown grid boxes have a reduction factor below 0.3.

384 As seen from (26), the intra-bin variance $\sigma_{\mathcal{B}o}^2$ of ship observations consists of sam-
 385 pling and measurement error variance components. Using (23), (28), and pooled esti-
 386 mates like (35), these components can be estimated separately; with averaging analo-
 387 gous to (41), obtain

$$388 \quad \hat{v}_o^{*2} = \sum_{m=1}^M \mu_m \sum_{y=1}^{Y_m} \gamma(y, m) [\mathcal{S}_{ao}(y, m)^2 + \mathcal{S}_{eao}(y, m)^2], \quad (42)$$

$$389 \quad \hat{\sigma}_o^{*2} = \sum_{m=1}^M \mu_m \sum_{y=1}^{Y_m} \gamma(y, m) [\mathcal{S}_d(y, m)^2 - \mathcal{S}_{eao}(y, m)^2], \quad (43)$$

390 where γ is defined by (36),(37). The intra-bin sampling \hat{v}_o^* and measurement $\hat{\sigma}_o^*$ ESDE
 391 for ship observations, computed by (42) and (43) are shown in Figures 3e,f. As with $\hat{\sigma}_{\mathcal{B}o}^*$,
 392 these are essentially ESDE components for a single observation, which are reduced by
 393 the factor $1/\sqrt{\mathcal{N}^*}$ (Figure 3d), when more observations are available.

394 5 Discussion

395 5.1 Sampling error

396 Estimate \hat{v}_o^* , given by (42) is based on the match-ups of the CCI Analysis SST and
 397 its uncertainty to the ship SST observations, a relatively small data sample. An estimate,
 398 based on the equation (21) that uses full CCI Analysis and its uncertainty

$$399 \quad \hat{v}_o^{*2} = \frac{1}{228} \sum_{m=1}^{12} \sum_{y=1}^{19} \frac{\mathcal{N}_a(y, m) - 1}{\mathcal{N}_a(y, m)} [\mathcal{S}_a(y, m)^2 + \mathcal{S}_{ea}(y, m)^2]$$

400 is shown in Figure 3g. Expectedly, this estimate is larger (by about 10% in areas of high
 401 DOF numbers) and smoother than the one based on the incomplete data (Figure 3e).
 402 It has the uncanny similarity in pattern, but generally is larger than the estimate pre-
 403 sented by Kennedy et al. (2011, their Figure 1d).

404 5.2 Measurement error

405 Kent and Challenor (2006) used the semivariogram method to estimate SST mea-
 406 surement error in 1970–1997 ICOADS data from ships. They identified pairs of ship SST
 407 observations made at the same hour and within 300 km of each other; squared differences
 408 between paired observations were binned by distance to construct the semivariogram;
 409 a linear fit to its points was extended towards zero distance separation to obtain the mea-
 410 surement error variance as the semivariogram’s nugget. Ship measurement ESDE in $30^\circ \times 30^\circ$
 411 averages from Kent and Challenor (2006, their Figure 2) is compared here with the mea-
 412 surement error estimates $\hat{\sigma}_o^*$, averaged to the same $30^\circ \times 30^\circ$ grid (Figure 3h,i). Two es-
 413 timates have a great deal of similarity (their pattern correlation is 0.75), despite the dif-
 414 ferences in the study period and estimation method. Relative difference ρ , shown in Fig-
 415 ure 3j has global RMS of 18.0%, with $|\rho| \leq 10\%$ in most of grid boxes. Grid boxes with
 416 $|\rho| > 10\%$ are generallyly in the areas of poor data coverage (cf. Figure 1d).

417 Kent and Berry (2008) introduced the measurement error model for marine observa-
 418 tions that combines random error with a “platform-dependent” bias or “micro-bias”,
 419 with the randomly distributed value over the platforms (ships). For this kind of error
 420 structure, if a bin contains many observations from a relatively small number of plat-
 421 forms, the error variance of its mean decreases inversely-proportionally to the number
 422 of platforms, rather than to the total number of observations. However, since moving
 423 ships, even at 14 knots (a relatively slow speed for modern ships), would cross the equa-
 424 torial $1^\circ \times 1^\circ$ bin in less than six hours (a typical time interval between ship observations),
 425 multiple observations from the same ship would not typically appear in the same bin,
 426 thus making equation (2) usable in this study.

427 Kennedy (2014, Table 1) listed published in 1965-2011 ship SST measurement ESDE
 428 that did not separate micro-biases from the purely random error parts. There are 19 es-
 429 timates there, ranging from 0.11°C to 3.5°C, with the median of 1.2°C, and 1–1.3°C in-
 430 terquartile range. Depending on the way of averaging measurement error estimates and
 431 especially on the averaging domain, global estimates can change appreciably. (Kent and
 432 Challenor (2006) report their global ESDE for ship SST random error as 1.2°C, if weighted
 433 by ocean area, and 1.3°C, if weighted by number of observations.) Estimates $\hat{\sigma}_o^*$ here
 434 can average to the global RMS of 1.14°C (Figure 3f), 1.13°C (Figure 3i), or 1.21°C, if
 435 the latter is constrained to the exact domain, where estimates in Figure 3h (global RMS
 436 of 1.26°C) are available.

437 6 Conclusions

438 Differences for 1992–2010 between monthly 1°×1° bin averages (for bins that con-
 439 tain more than one observation) of ICOADS ship SST and of the ESA SST CCI Anal-
 440 ysis are presented here as the sum of their climatological bias component and remain-
 441 ing residuals (anomalies) with magnitudes that agree with the random error model in
 442 the areas of sufficient data coverage. The model assumes for ship observations the i.i.d
 443 measurement and sampling errors within bins and high intra-bin correlation for the CCI
 444 Analysis uncertainty. Location-dependent estimates of ship SST measurement and sam-
 445 pling error were obtained. Estimates of sampling error are similar in pattern, but larger
 446 than those previously published. Ship SST measurement error is consistent with previ-
 447 ous estimates in spatial pattern and global RMS (1.13-1.21°C, depending on the aver-
 448 aging domain and procedure).

449 Acknowledgments

450 To the memory of M. Benno Blumenthal (1959-2018), creator of Data Library and In-
 451 grid software, with which all figures in this Letter were calculated and plotted. Discus-
 452 sions with P.C.Cornillon, E.C.Kent, C.J.Merchant, J.J.Kennedy, and N.A.Rayner are grate-
 453 fully acknowledged. ICOADS, Release 3.0, is obtained at

454 <https://rda.ucar.edu/datasets/ds548.0/>

455 ESA SST CCI Analysis for 9/1991–12/2010, version 1.0, is obtained at

456 <https://catalogue.ceda.ac.uk/uuid/916986a220e6bad55411d9407ade347c>

457 Supported by grants OCE-1853717 from NSF and NA17OAR4310156 from NOAA. LDEO
 458 contribution number xxxx.

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